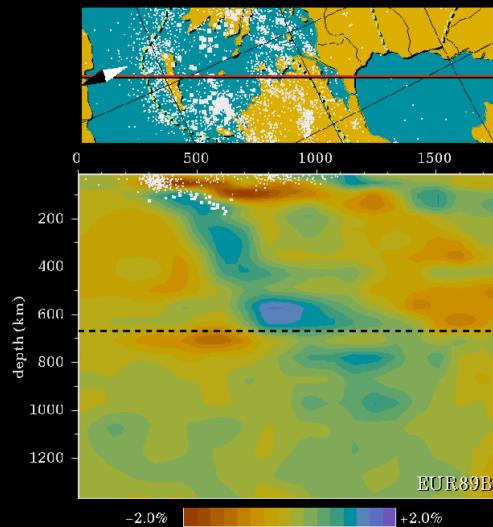


# Seismic Tomography



mini course  
*Earthscope workshop*  
*Monterey CA, 2007*  
**Suzan van der Lee**

Examples for  
North America



## Seismic Tomography

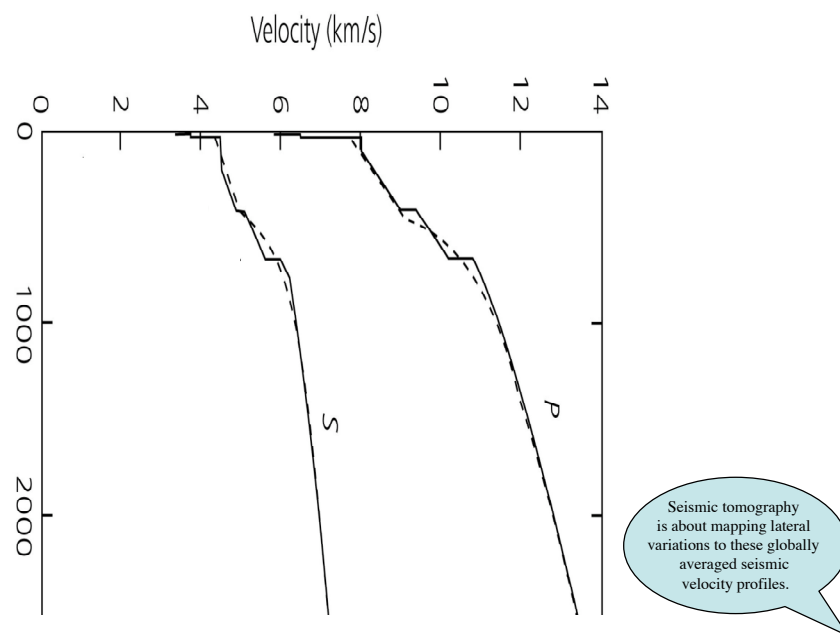
Mini-Course with contributions and inspiration from:

*John VanDecar,  
Christian Schmid,  
Guust Nolet,  
Steve Grand,  
Saskia Goes,  
Ken Dueker  
Heather Bedle,  
MOMA array research team,  
Berkeley Seismic Group*

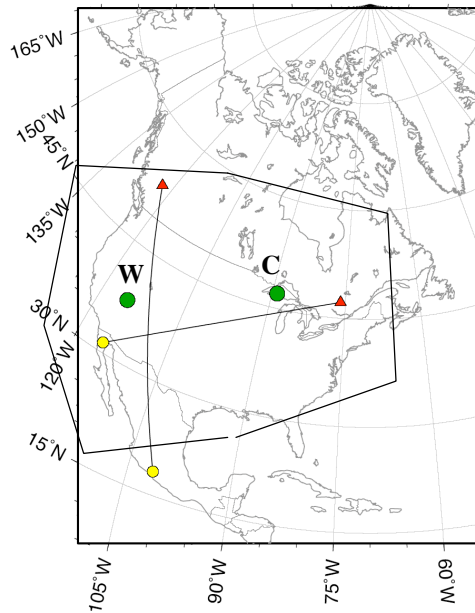
# Seismic Tomography

1. A Simple Exercise
2. Some Nitty Gritty
3. Resolution
4. Model Comparisons
5. Tips

## Seismic Tomography: A Simple Exercise



## Seismic Tomography: A Simple Exercise



consider two regions:

1. **W**estern North America
2. **C**entral North America

consider two seismograms:

1. Gulf of CA to Quebec:  
waves on time
2. Cocos trench to BC:  
waves late

Seismic waves penetrate deeply in the Earth's mantle and crust. The horizontal propagation velocity of seismic waves depends on the S-wave velocity with depth, as well as on the frequency and tone of the waves. In this simplified minicourse we incorrectly assume that the horizontal propagation velocity equals the S-velocity of the uppermost mantle.

## Seismic Tomography: A Simple Exercise

1. distance  $x = 3600$  km  $dt = 0$  s  $x_W = x_C = 1800$  km
2. distance  $x = 4000$  km  $dt = 52$  s  $x_W = x$

Reference slowness  $s_0 = 0.222$  s/km  $= 1/(4.5$  km/s)

$x = vt \rightarrow x/v = t \rightarrow xs = t$ , with  $s = 1/v$

distance\*slowness = time  $\rightarrow$

distance\*slowness difference ( $ds$ ) = time difference ( $dt$ )

Two independent measurements ( $dt_1$  and  $dt_2$ ) yield two equations to be solved for two unknowns ( $ds_W$  and  $ds_C$ ):

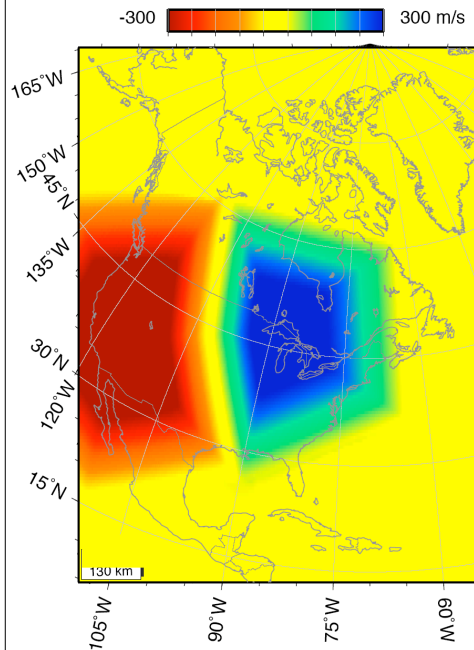
$$1. \quad 1800ds_W + 1800ds_C = 0$$

$$2. \quad 4000ds_W = 52$$

Solution:  $ds_W =$  s/km

$ds_C =$  s/km

## Seismic Tomography: A Simple Exercise



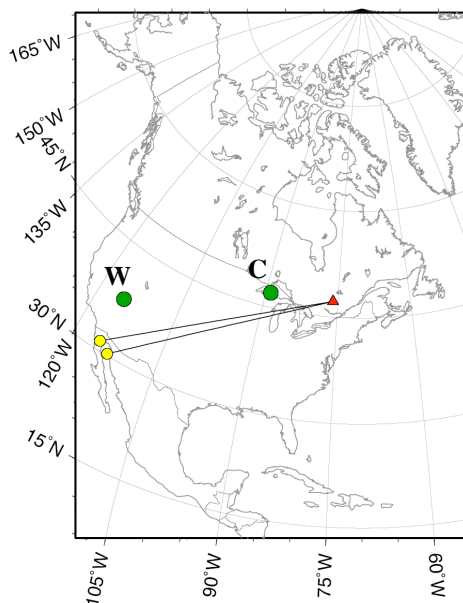
solve 2 equations for 2 unknowns:

1. central North America "is" fast
2. western North America "is" slow

1. The stable Precambrian lithosphere of central North America is cool and rigid, allowing seismic waves to propagate efficiently and rapidly.
2. The mantle beneath Phanerozoic western North America is hot and weak, hindering the efficient propagation of seismic waves, slowing them down.

Our model has only two variables: the velocity difference for West and Central North America. Everywhere else we keep the velocity difference fixed to zero (yellow).

## Seismic Tomography: Some Nitty Gritty



Same two regions (**W** and **C**)

consider 2 *different* seismograms:

1. Gulf of CA to Quebec: waves on time
2. Baja to Quebec: waves almost on time

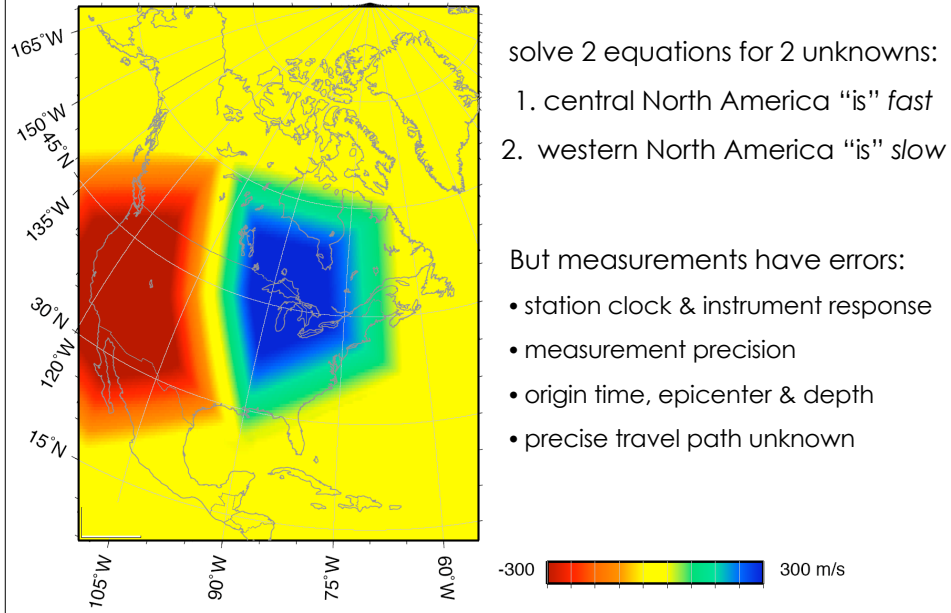
Again: two eqns & two unknowns:

1.  $1800ds_W + 1800ds_C = 0$
2.  $1900ds_W + 1800ds_C = 1.3$

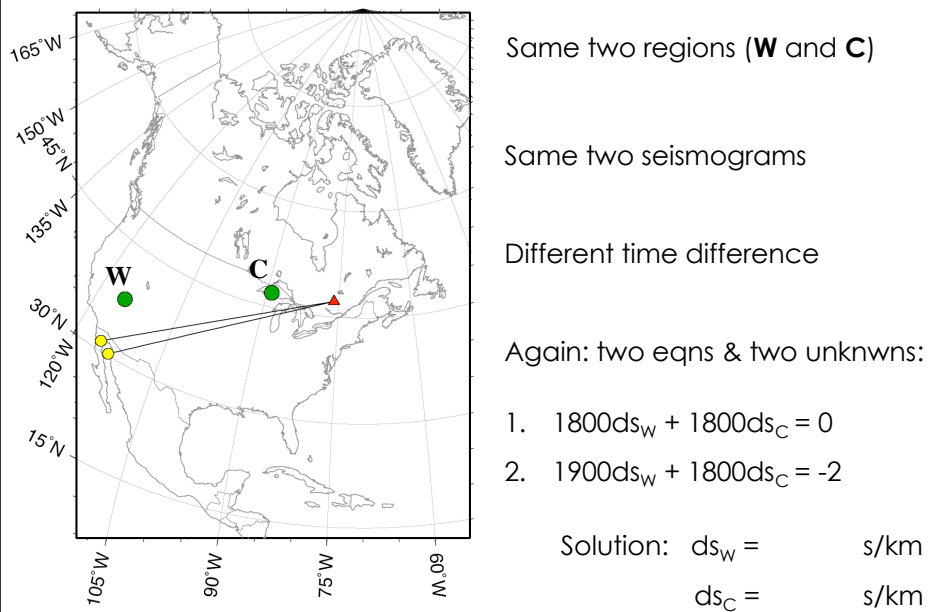
Solution:  $ds_W =$  s/km  
 $ds_C =$  s/km



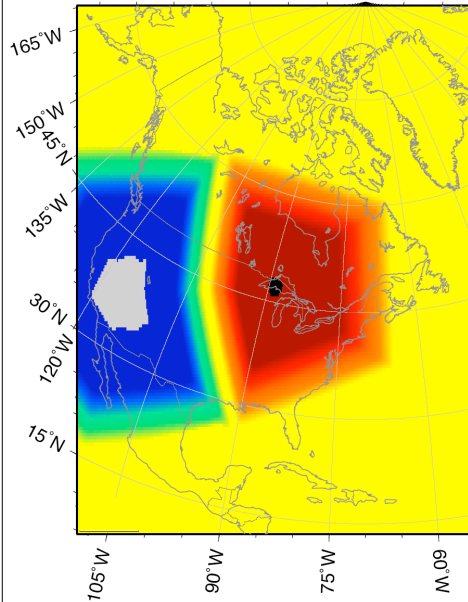
## Seismic Tomography: Some Nitty Gritty



## Seismic Tomography: Some Nitty Gritty



## Seismic Tomography: Some Nitty Gritty



solve 2 equations for 2 unknowns:

1. central North America "is" **slow**
2. western North America "is" **fast**

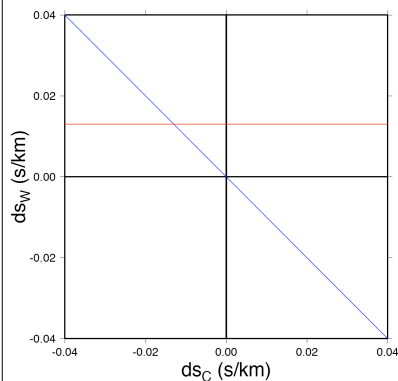
??

What went wrong?



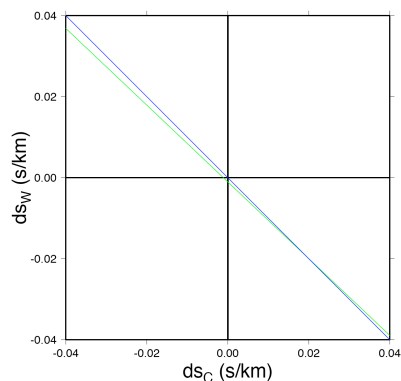
The grey and black areas indicate that the velocity difference is "off the chart".

## Seismic Tomography: Some Nitty Gritty



Case 1: The 2 seismograms clearly identify one crossover point in model space: (-0.013, 0.013)

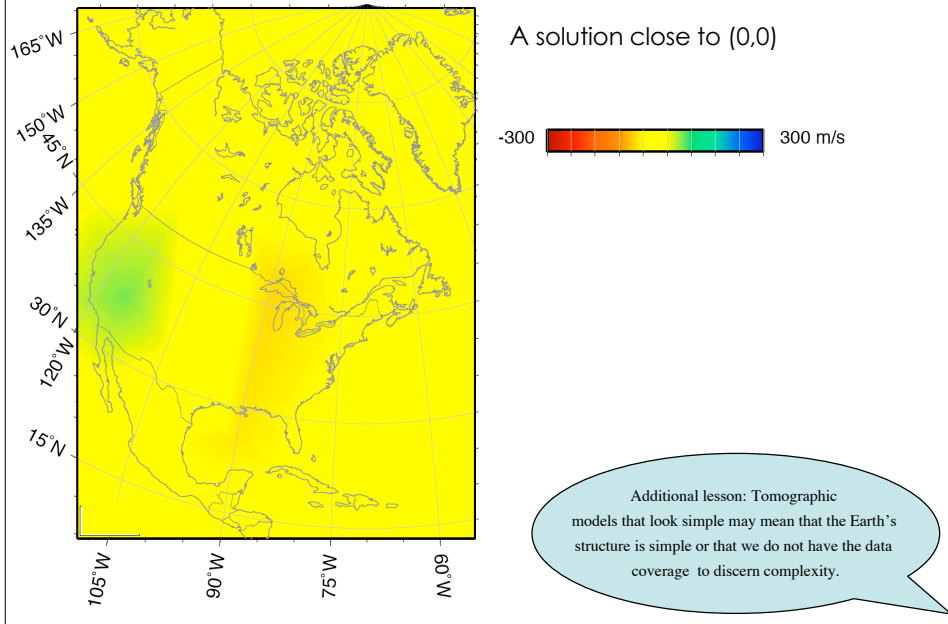
Case 2 and milder variants are common in seismic tomography, at least for a subset of model variables. Case 1 is also common, but unfortunately rarely for more than a subset of model variables.



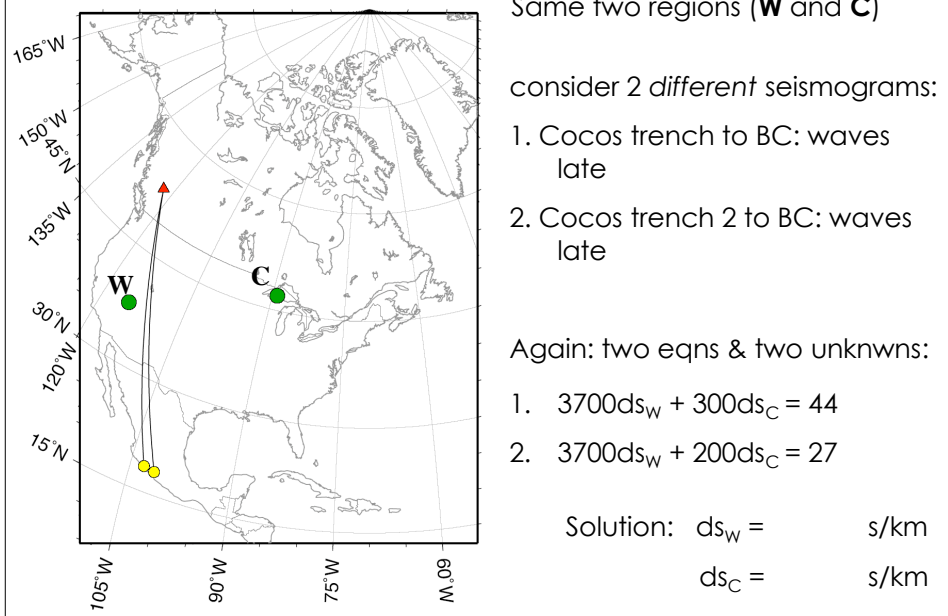
Case 2: The 2 seismograms are so close that, within error, the crossover point is no more meaningful as a solution than any other point on the 2 lines, including: (-0.013, 0.013)

→ New strategy: Pick a solution close to (0,0)

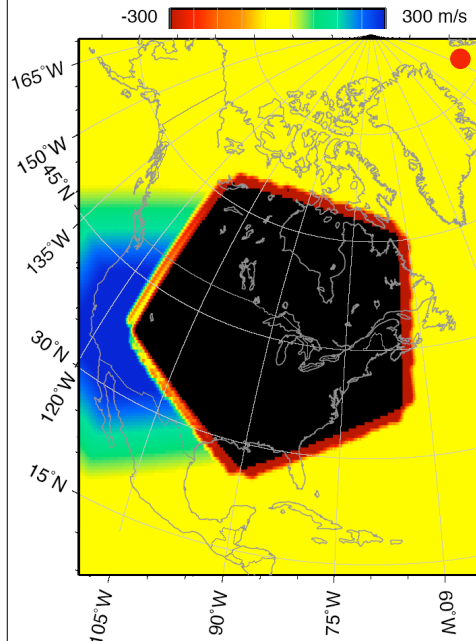
## Seismic Tomography: Some Nitty Gritty



## Seismic Tomography: Some Nitty Gritty



## Seismic Tomography: Some Nitty Gritty



solve 2 equations for 2 unknowns:

1. central North America "is" **slow**
2. western North America "is" **fast**

??

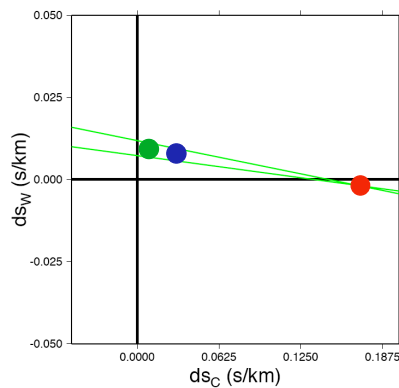
Apply new strategy:



pick a solution close to (0, 0)

Again, small error in the data lead to large errors in the imaged velocity differences. The negative velocity difference for Central North America is, again, off the chart.

## Seismic Tomography: Some Nitty Gritty

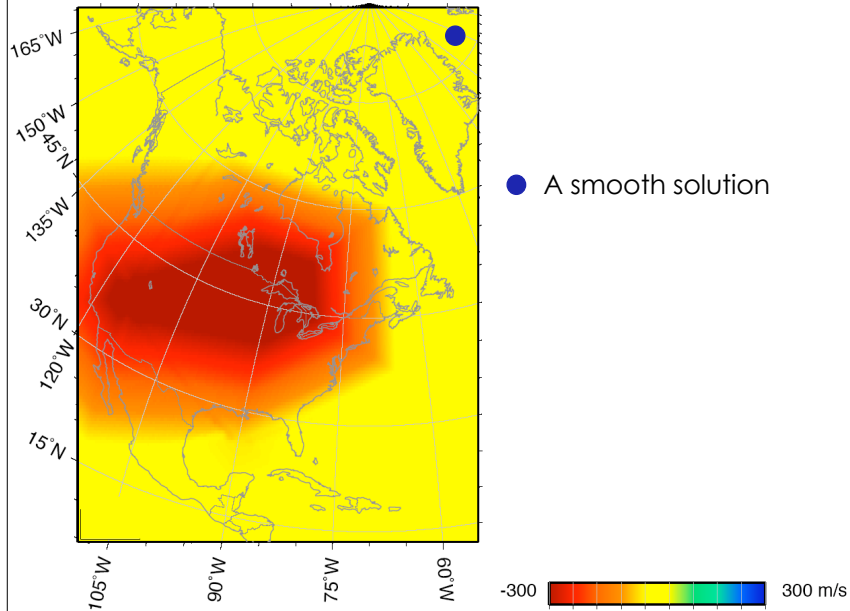


- formal solution (previous map)
- solution close to (0, 0)
- smooth solution (next map)

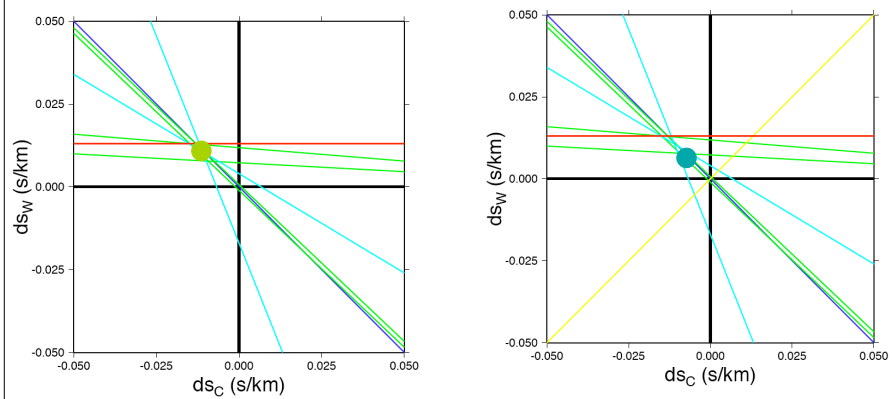
New strategy:

pick a smooth solution:  
values are like each other

## Seismic Tomography: Some Nitty Gritty



## Seismic Tomography: Some Nitty Gritty

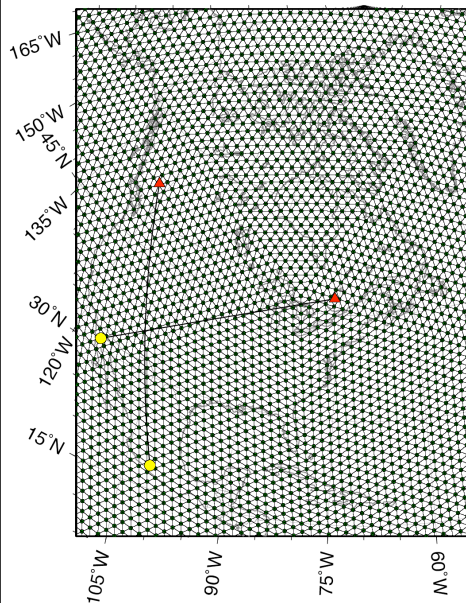


- least-squares solution = the solution that is closest to all lines

The lines shown are from the previous examples; a set of equations with random errors. "Damping" will draw the solution towards the origin. "Smoothing", or technically "flattening", will draw the solution towards the yellow line.

- damped/smoothed least-squares solution → the solution underestimates actual slowness values

## Seismic Tomography: Some Nitty Gritty



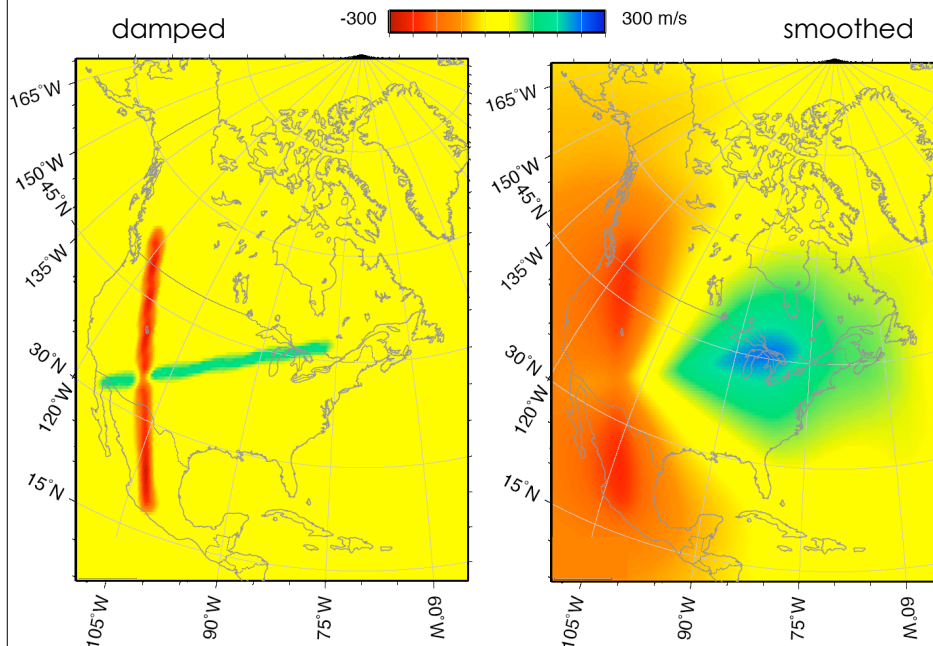
Many more than 2 regions

Two seismograms

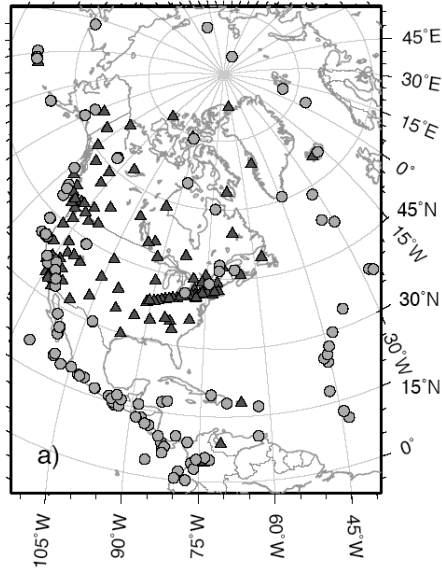
→ damping/smoothing

Beyond west and central North America:  
Now consider many more than two model variables,  
but still with only two equations that  
constrain the model variables.

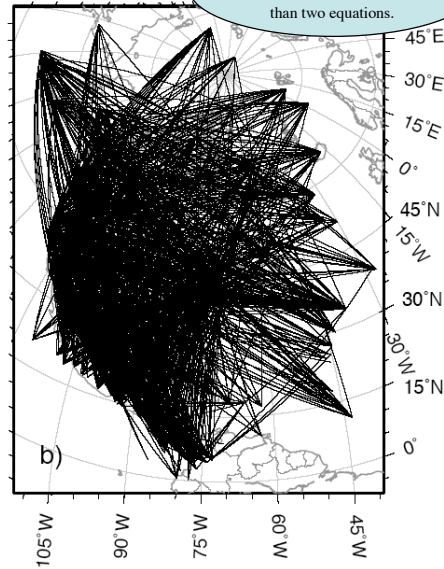
## Seismic Tomography: Some Nitty Gritty



## Seismic Tomography: Some Nitty Gritty

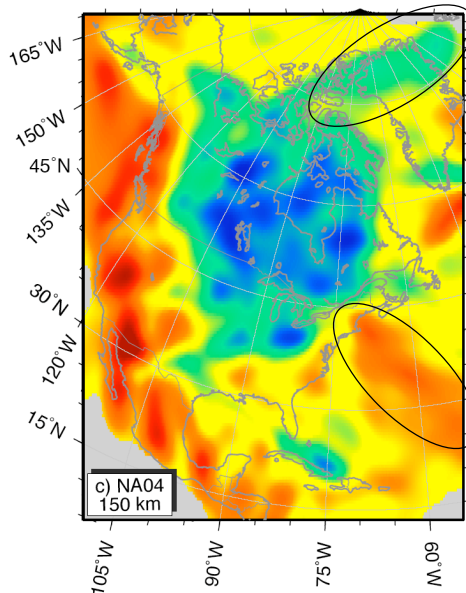


Many more than 2 seismograms



Note difference in coverage between continent and ocean

## Seismic Tomography: Some Nitty Gritty

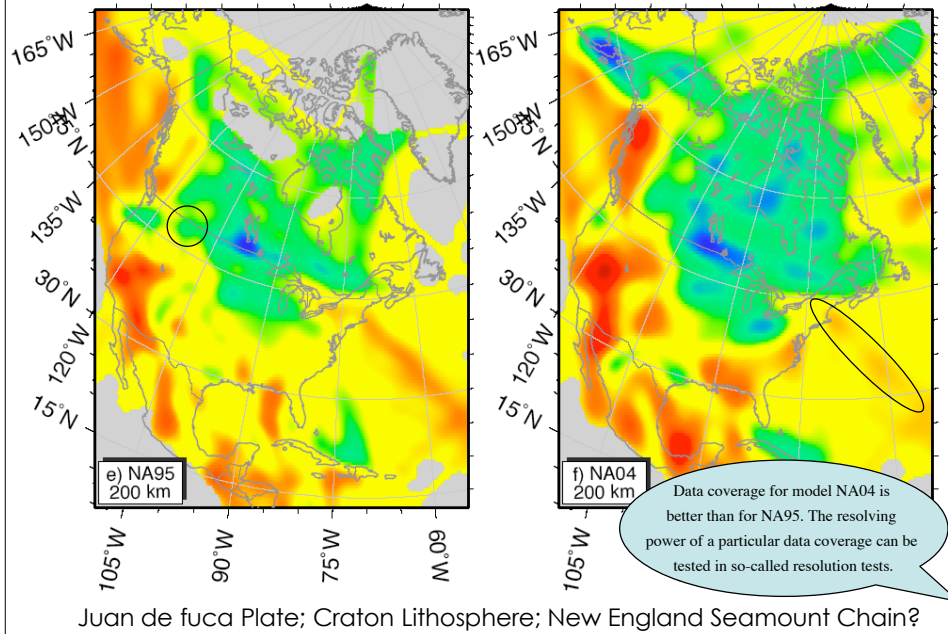


Example 1: N Greenland: structure **smoothed** along wave paths from continent into ocean (even though *true structure* may have edge at margin), but **not smoothed** much in N-S direction (*true structure* may not be so fast in better covered South)

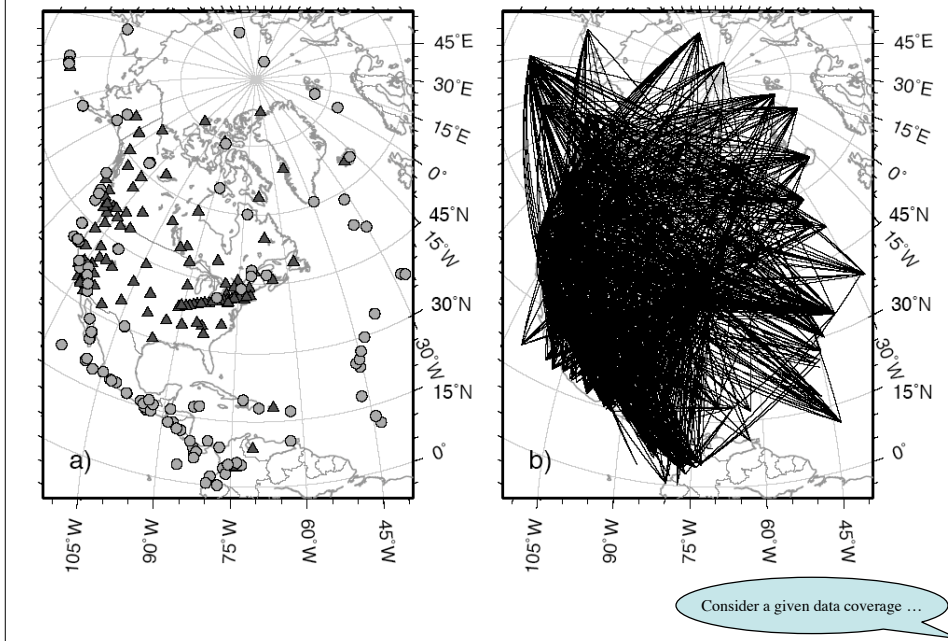
Example 2: Atlantic Ocean: structure **smoothed** along wave paths between ridge and continental margin (even though *true structure* may change with age of ocean floor), but **not smoothed** much parallel to the ridge/margin (even though in that direction *true structure* may be more homogeneous)



## Seismic Tomography: Some Nitty Gritty



## Seismic Tomography: Resolution

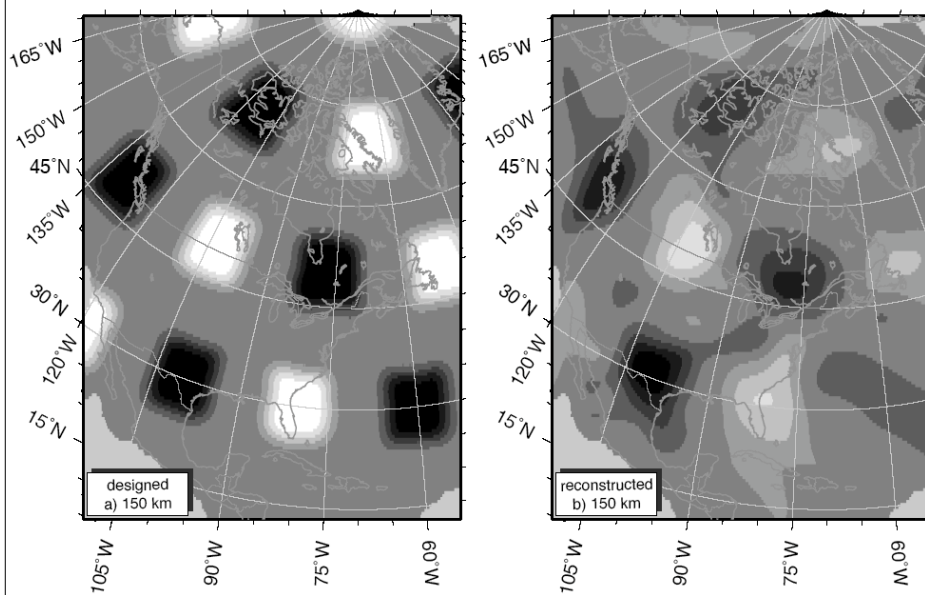




...then consider a hypothetical design of velocity differences, compute data for these differences, add random noise to these "data", and use the "data" to reconstruct the model (next slide).

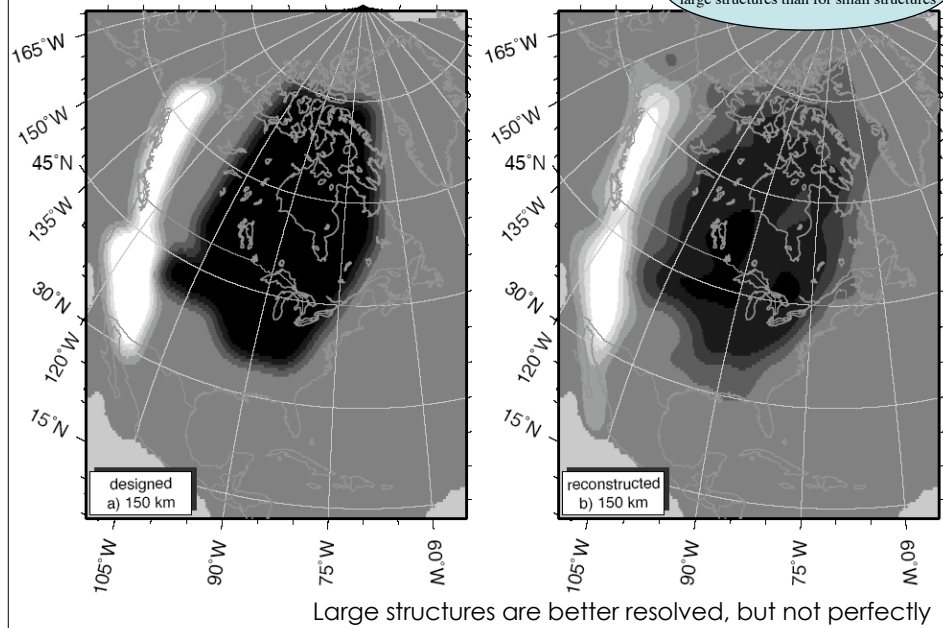
Model-specific lessons (next slide): In the Atlantic Ocean our data cannot resolve the difference between a single checker or a chain of anomalous velocity. Better data coverage does a better job at defining the edge of the North American Craton, particularly in the presence of nearby similarly rigid structures with a different tectonic origin (such as subducted lithosphere or old oceanic lithosphere).

## Seismic Tomography: Resolution

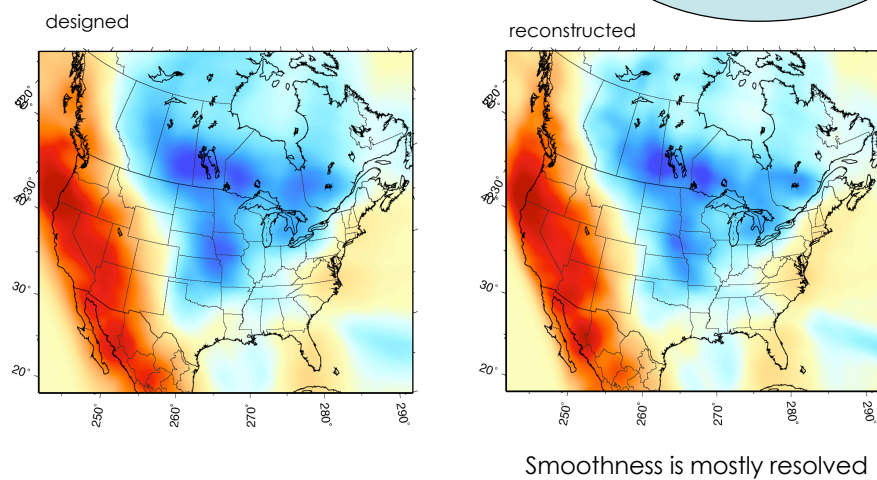


Velocities are underestimated, and heterogeneously resolved

## Seismic Tomography: Resolution

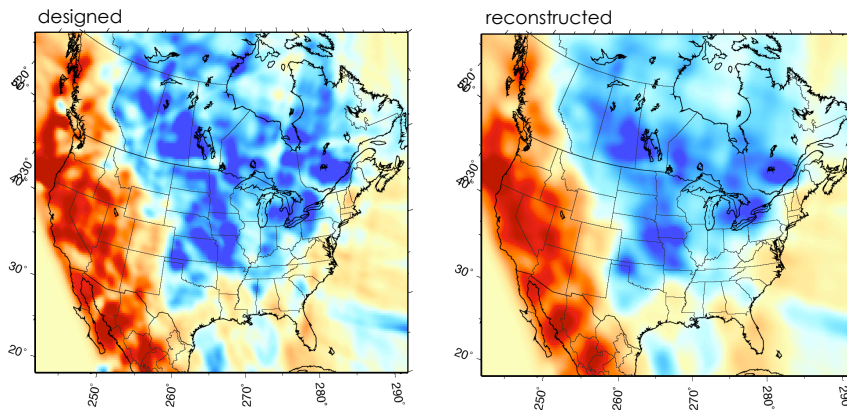


## Seismic Tomography: Resolution



## Seismic Tomography: Resolution

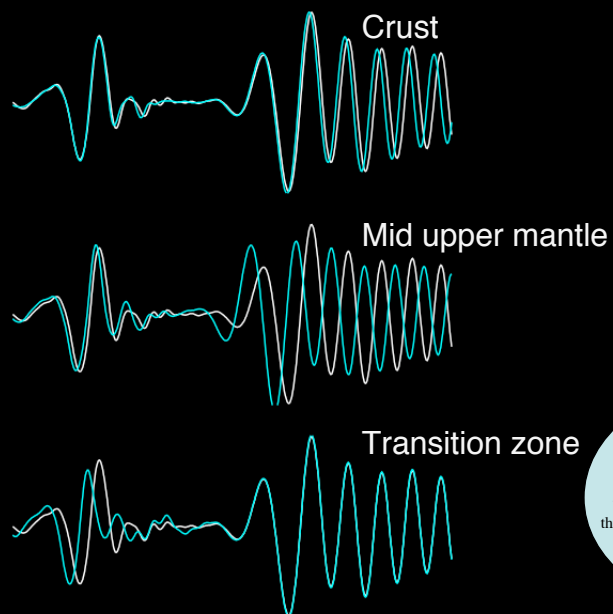
...as well as rough structures.  
Our damping/smoothing choices filter  
out some small-scale structures while  
having the power to resolve  
smooth structures.



Roughness is also resolved, but not entirely

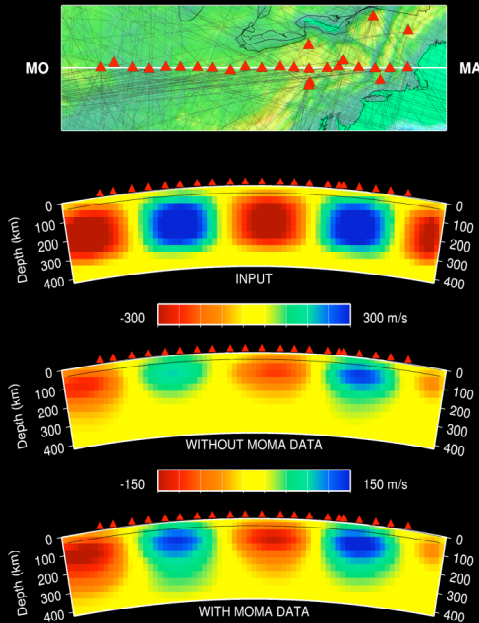
## Seismic Tomography: Resolution

Depth resolution  
comes from  
different waves  
(tones) and  
frequencies



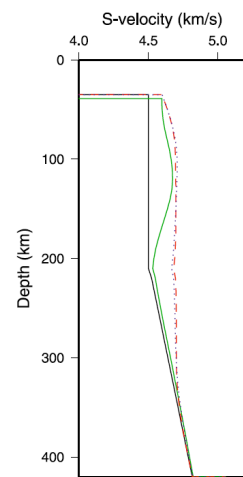
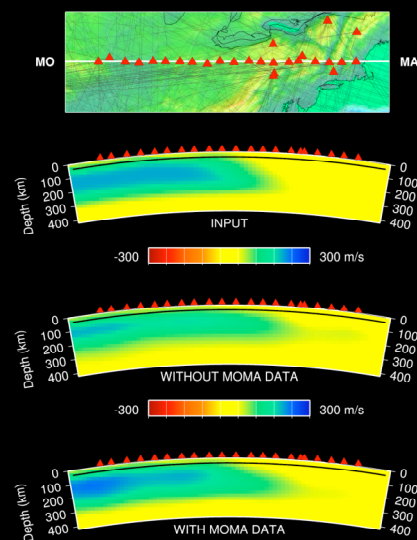
The three turquoise  
synthetic seismograms are  
for three different models with  
a velocity anomaly in one of each of  
three depth regions. They are compared  
with the same white reference  
synthetic seismogram.

## Seismic Tomography: Resolution



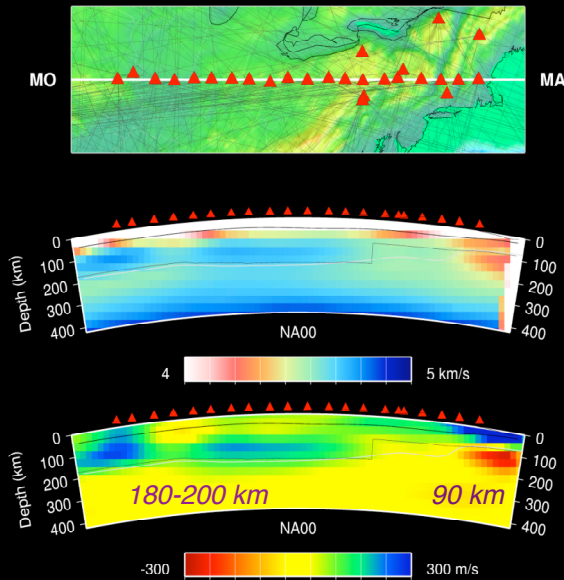
Like **USArray's** Transportable Array, the dense MOMA station line improves resolving power of the data set, both in depth and for absolute values of velocity heterogeneity.

## Seismic Tomography: Resolution



Improved sub-lithospheric resolution by incorporating waveform data from dense seismic arrays (e.g. MOMA)

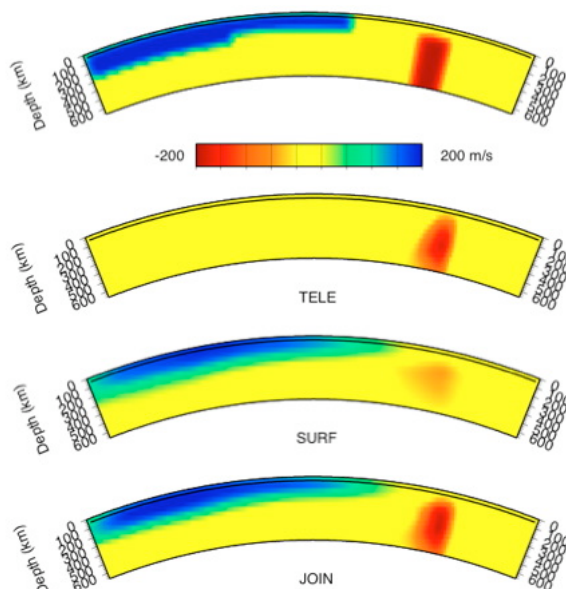
## Seismic Tomography: Resolution



Well-resolved seismic thickness estimate agrees with

- petrologic thickness from xenoliths
- thermal thickness from mineral physics and tomography
- thermal thickness from heat flow
- twice the elastic thickness from topography and gravity

## Seismic Tomography: Resolution



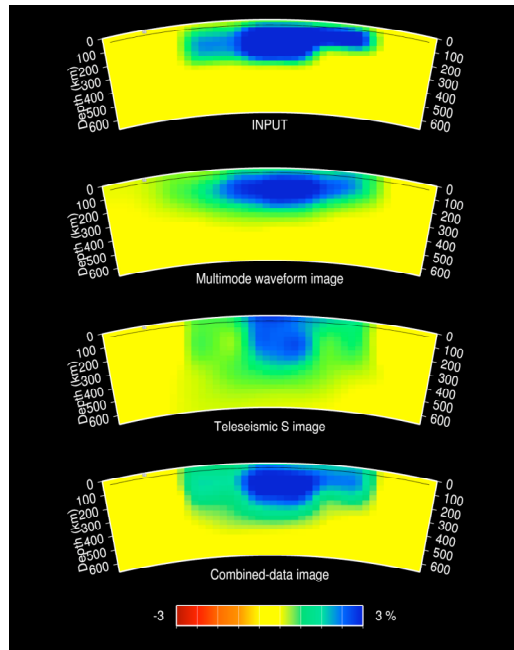
A. Teleseismic arrivals sense structures directly beneath stations

B. Regional waveforms sense structures between stations

A. Teleseismic arrivals sense structures relative to unknown average

B. Regional waveforms sense absolute structures

## Seismic Tomography: Resolution

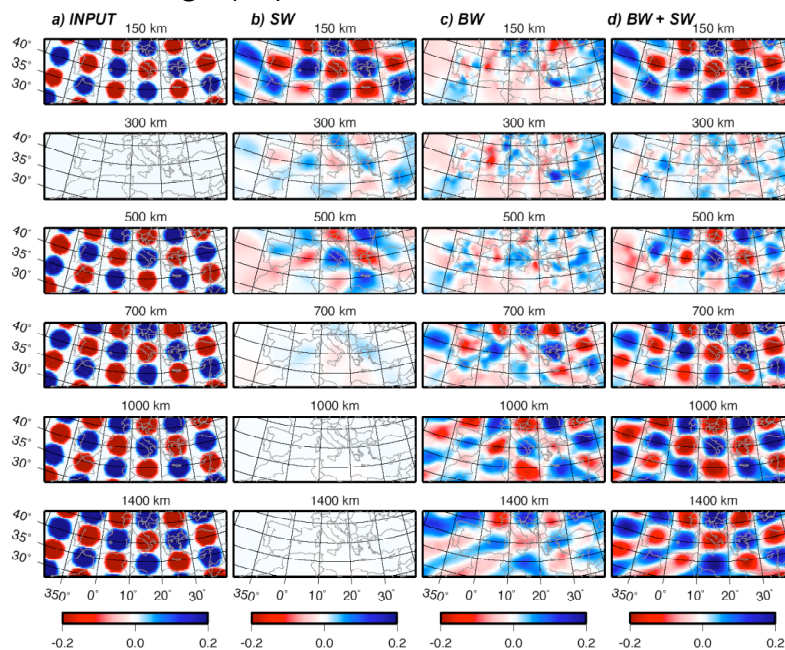


A. Teleseismic arrivals sense structures with greater lateral resolution

B. Regional waveforms sense structures with greater depth resolution

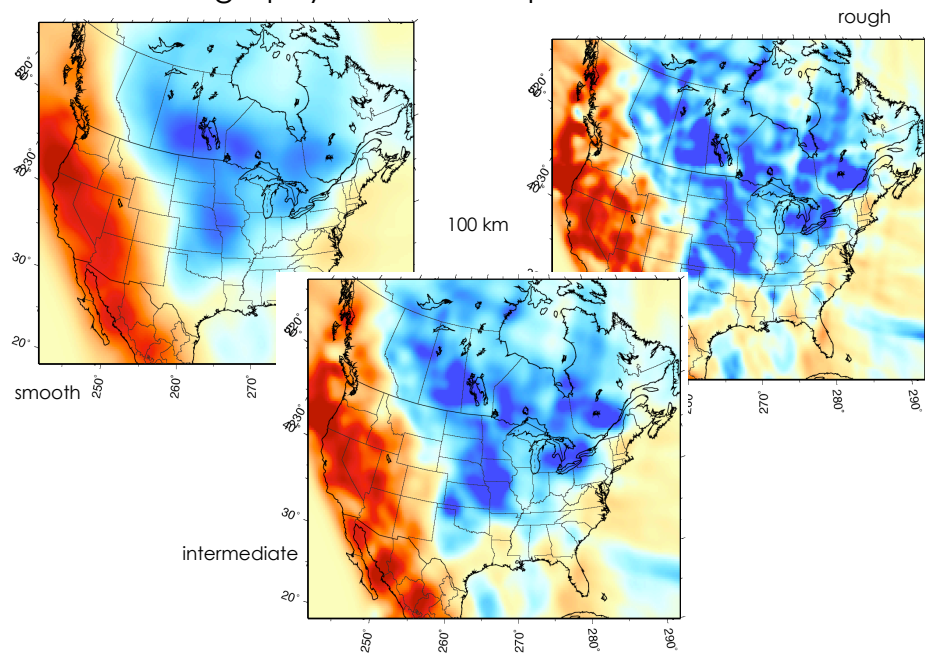
Combined tomography is able to discriminate between a thick, normal and a thin, strong lithosphere

## Seismic Tomography: Resolution

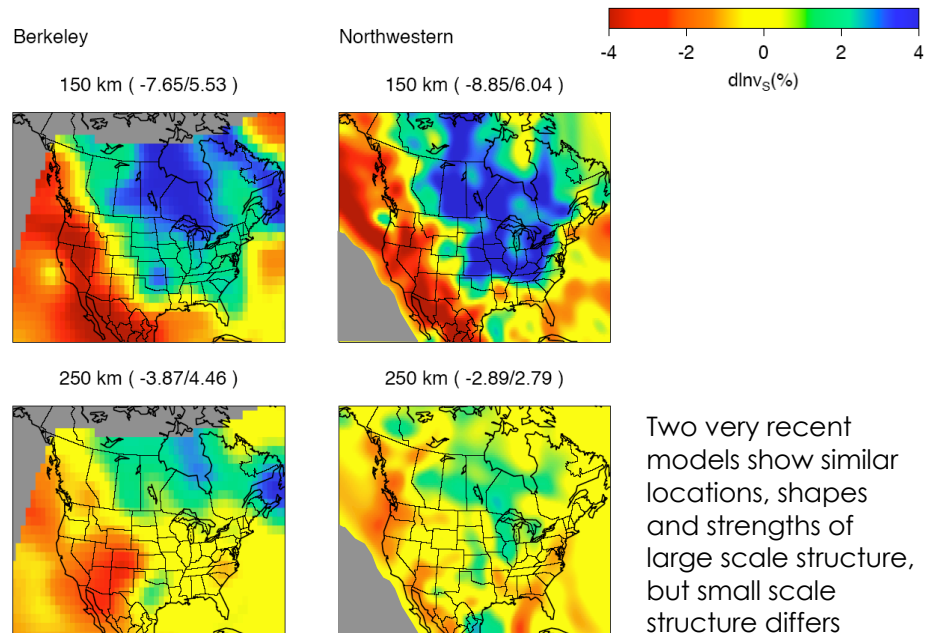




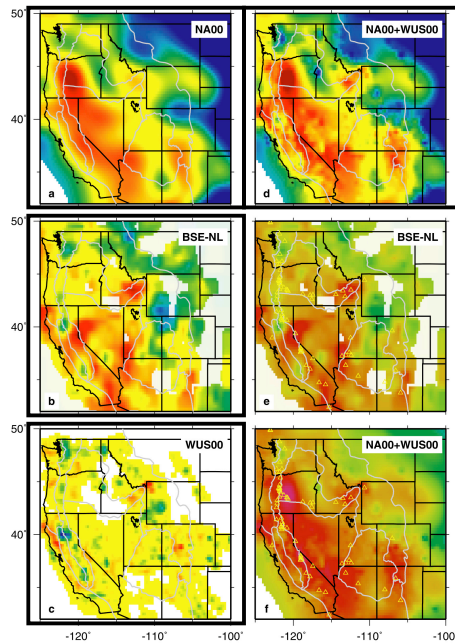
## Seismic Tomography: Model Comparisons



## Seismic Tomography: Model Comparisons

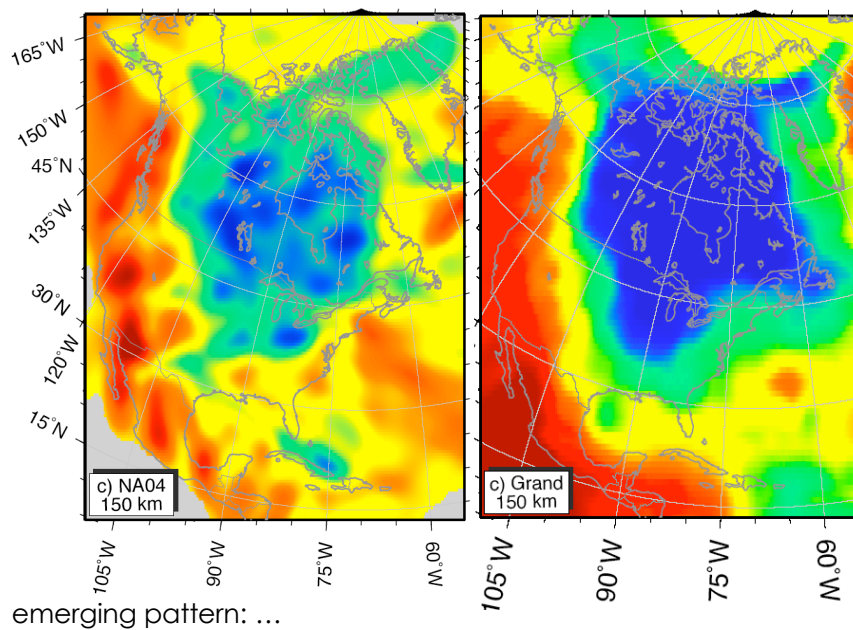


## Seismic Tomography: Model Comparisons



Two older models (one very smooth, one very rough) show practically nothing in common, but when combined are similar to third, intermediate-scale model

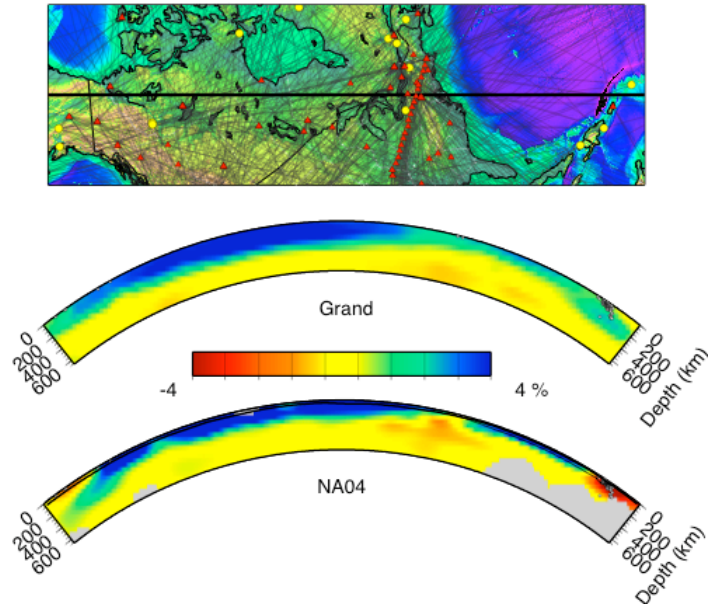
## Seismic Tomography: Model Comparisons



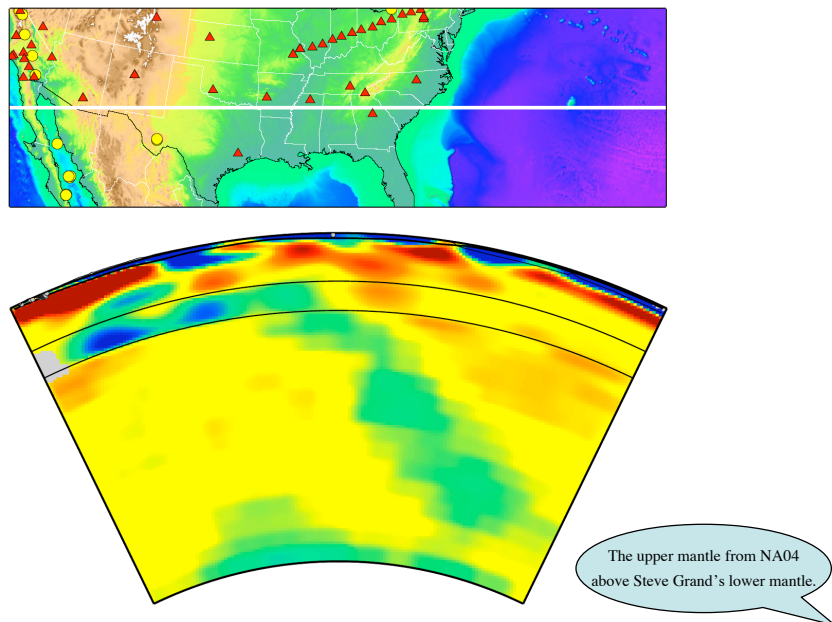
emerging pattern: ...



## Seismic Tomography: Model Comparisons



## Seismic Tomography: Model Comparisons



## Seismic Tomography: Tips

*When seriously interested in a tomographic model:*

1. **Read** the associated peer-reviewed publication
2. Examine the accompanying **resolution test** results
3. **Compare** the model with other models for the region
4. **Discuss** the model and its details with your seismology colleagues and or its authors

The End